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Abstract

- - strategies.
 - cloud brightening.
- "The effect of anthropogenic aerosols on cloud droplet the industrial period." [Carslaw, Nature, 2013]
- aerosols behave at fine scales.
 - transport/mixing/spreading.

Mathematical Model We consider a conservative tracer transported via Our climate is changing—this may get unpleasant. turbulent atmospheric flow according to the advection-**Global climate models are:** diffusion equation Critical to understanding these changes. The best option to study dramatic intervention E.g., stratospheric aerosol injection and/or marine The tracer mass is divided among particles with (a.s.) Known to be imperfect. distinct positions and masses concentrations and radiative properties is the source of one of the largest uncertainties in the radiative forcing of climate over such that the concentration field may be recovered as **Goal:** Gain a better understanding of how injections of $q(\boldsymbol{x},t) = \int_{\mathbb{R}^d} \sum_{i=1}^d m_j \phi(\boldsymbol{x}-\boldsymbol{z}) \delta(\boldsymbol{z}-\boldsymbol{x}_j) d\boldsymbol{z}$ Characterize the sub-grid dynamics of aerosol plume **Ok, sounds good—How?** Lagrangian (particle) model **d** Data sources: LES results (Blossey, Wood, McMichael—UW) Fog chambers (Pattyn, Zenker, Wright, Sanchez) (Schmidt, Patel, Bosler, UW Team) Stochastic wind parameterization. Satellite imagery of ship tracks

- for aerosol transport.
- Resolves sub-grid turbulent effects that drive plume spreading.

- Temporary cloud trails from ship exhaust cloud seeding.
- Actual example of anthropogenic MCB. Apply ML model to generate parameterizations. (Patel, Shand, Shuler, Warburton)
- Automated image recognition and feature extraction.





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The particles move under the influence of wind, and particle positions evolve according to

 $\coloneqq \sum m_j \delta(\boldsymbol{x} - \boldsymbol{x}_j).$

 $=\sum m_j \phi(\boldsymbol{x}-\boldsymbol{x}_j)$

$$\frac{d\boldsymbol{x}_i}{dt} = \boldsymbol{v}_i^*(t).$$

 $\frac{\partial q}{\partial t} = -\nabla \cdot (q\boldsymbol{v}), \qquad q(\boldsymbol{x},0) = g(0; \boldsymbol{\mu} = \boldsymbol{x}_0, \boldsymbol{\sigma}_0),$

 $x_i(t), i = 1, ..., N_p, \qquad m_{\text{TOT}} = \sum m_j(t).$

 $v(x,t) = f(x,t), \quad v(x,0) = f(x,0),$

 $x \in \mathbb{R}^d, \quad t > 0.$

We impose a stochastic parameterization for wind velocity

$$\mathrm{d}\boldsymbol{v}_{i}^{*}(t) = \frac{\overline{\boldsymbol{v}} - \boldsymbol{v}_{i}^{*}(t)}{T_{\ell}} \mathrm{d}t + \sqrt{\frac{2\boldsymbol{\eta}}{T_{\ell}}} \mathrm{d}\boldsymbol{W}(t),$$

and integrate in time as

$$\boldsymbol{v}_i^*(t + \Delta t) = \boldsymbol{v}_i^*(t) + \frac{\left[\overline{\boldsymbol{v}} - \boldsymbol{v}_i^*(t)\right]\Delta t}{T_\ell} + \sqrt{\frac{2\boldsymbol{\eta}\Delta t}{T_\ell}} \mathrm{d}\boldsymbol{V}$$



OF ATMOSPHERIC AEROSOL TRANSPORT



